

Recent developments inside a Lagrangian Particle Dispersion Model for the estimation of concentration peaks

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Structure of the presentation

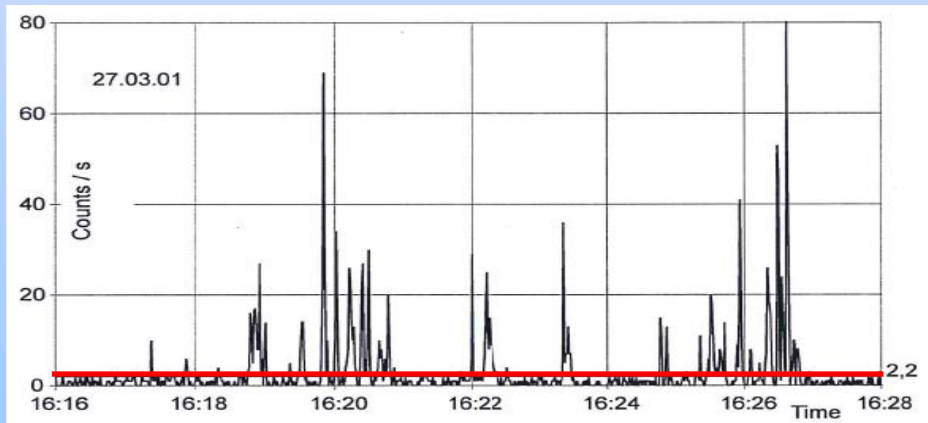
1. General description of the problem
2. Peak-to-mean formulations
 - a) *Longitudinal*
 - b) *Variance-transport*
3. Model implementations into the SPRAY LPDM
4. Results in controlled/real conditions and comparisons between different methods
5. Conclusions



1. General description of the problem

Some fundamental remarks

- Olfactory problems are perceived by an individual through a single respiratory act, lasting about 5 seconds
- it is the concentration averaged during this brief sampling period that does matter in principle, and not the relatively long-term average concentration (eg one hour)



Instantaneous (—) vs
average (—) concentrations

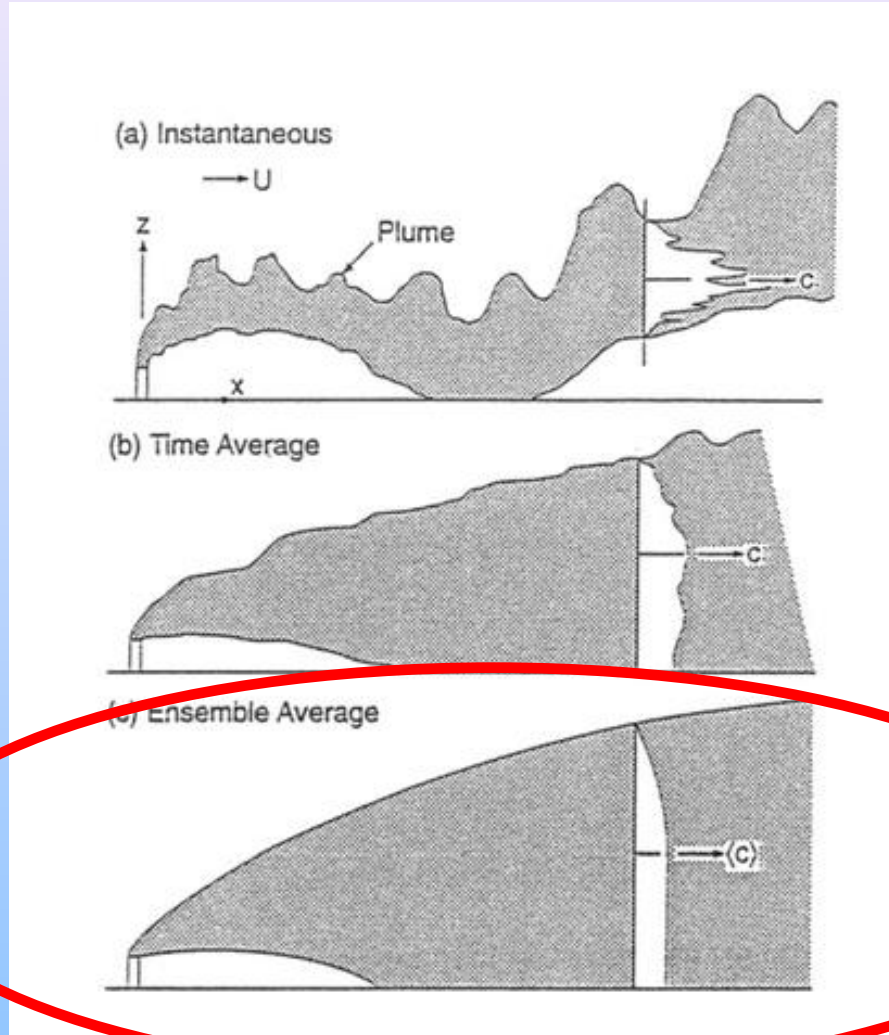
⁸⁵Kr emitted in air (Lung e al. 2002)

It is therefore necessary to model the "instantaneout" values (or at least at a frequency of the order of 5 seconds) of the concentration



1. General description of the problem

Various way to compute the concentrations



This is what is typically done by dispersion models !!!



2. Peak-to-mean formulations



The method consists in estimating the peak (short time/instantaneous) concentration values through the average values calculated by the model, using a multiplication factor (Peak-to-mean factor)

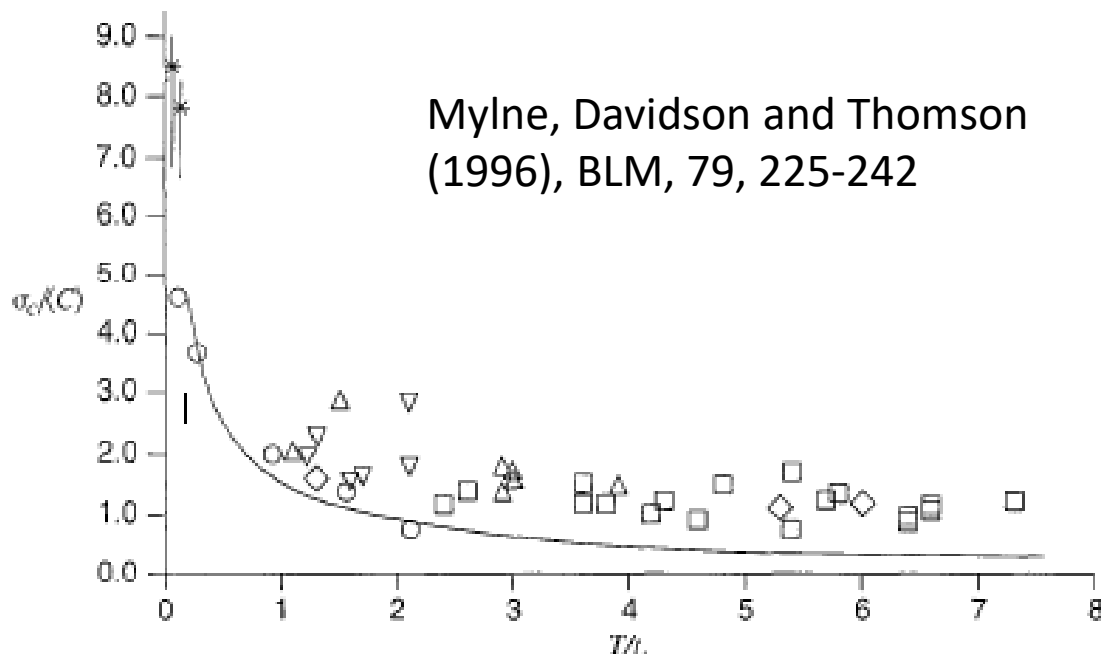
Some reference guidelines provides for the use of a **constant Peak-to-mean factor F** (independent from space, time, met)

For example: $F=4$ (Germany), $F=2.3$ (Italy)



2. Peak-to-mean formulations

a) Longitudinal Peak-to-mean Ratio



Starting from observations of the concentration variance inside the plume as a function of the flying time T , the method provides an asymptotic $p/m=1$ for large T and larger values for small T

$$p/m = \Psi(x) = 1 + (\Psi_0 - 1) \cdot \exp\left[-0.732 \cdot \frac{x}{U \cdot T_L}\right]$$

where

$$\frac{C_p}{C_m} = \left(\frac{t_m}{t_p}\right)^\alpha = \Psi_0$$

(Piringer et al., Atmos. Env., 2017)

It is assumed that the p/m at the source position is a function of the atmospheric stability α , the typical average time t_m (3600 s) and to the average time t_p over a short period (e.g. 5 s)



2. Peak-to-mean formulations

b) Simplified Variance-Transport equation

Considering $C = \langle C \rangle + c'$ and starting from the general transport equation for c'^2

$$\frac{\overline{\partial c'^2}}{\partial t} + U_j \frac{\overline{\partial c'^2}}{\partial x_j} = -2\nu_c \overline{\left(\frac{\partial c'}{\partial x_j}\right)^2} - \frac{\partial}{\partial x_j} \left(\overline{u'_j c'^2} \right) - 2\overline{u'_j c'} \frac{\partial \langle C \rangle}{\partial x_j}$$

Following Oetli and Ferrero (2017): A simple model to assess odour hour for regulatory purposes, *Atmospheric Environment*, G155, **162-173**

- neglecting the transport terms
- neglecting the diffusion term
- comparing the rest to the lagrangian equations

the Variance-Transport equation is reduced to:

$$\frac{\overline{\partial c'^2}}{\partial t} = -2\sigma_{ui} T_{Li} \left(\frac{\partial \langle C \rangle}{\partial x_j} \right)^2 - \frac{\overline{c'^2}}{2T_{Lw}} \quad \text{having the analytical solution:} \quad \overline{c'^2} = 2\sigma_{ui} T_{Li} (2T_{Lw}) \left(\frac{\partial \langle C \rangle}{\partial x_j} \right)^2$$

the p/m is estimated from $\langle C \rangle$, $\overline{c'^2}$ computed at each point and a PDF of a given form



The SPRAY-3 Lagrangian Particle Dispersion Model is considered

Longitudinal P/M

- a value $\Psi(p)$ is associated to each computational particle
- the previous equation is applied in differential form, considering $\Delta x = U \Delta t$
- each particle contributes with its own $\Psi(p)$ inside a concentration cell to the computation of an averaged Peak-to-mean factor
- U and T_L are space/time dependent, computed on each particle position

Variance-transport P/M

- each 'sampling Δt ' a $\frac{\partial \langle C \rangle}{\partial x_j}$ field is computed on an Eulerian grid
- the simplified equation is applied to compute $\overline{c'^2}$
- using $\overline{c'^2}$ and $\langle C \rangle$, a p/m is computed using a Weibull PDF (other options are possible)
- the peak concentration values are estimated as the 98 percentile of the Weibull PDF with the given $\overline{c'^2}$ and $\langle C \rangle$



Setup of the numerical experiments in simple controlled conditions

Homogeneous and stationary conditions, 12 x 12 km² domain, flat terrain, 1 and 4 m/s wind, 2 different stability conditions, point source

Point source characteristics

Stack height = 10 m

Diameter = 0.5 m

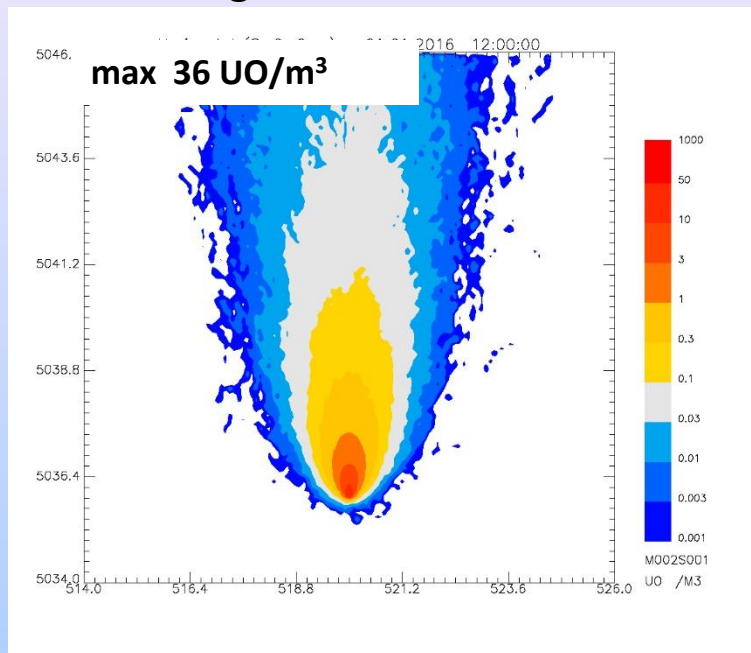
Exit temperature = environment

Emitting flow = 70000 UO/s



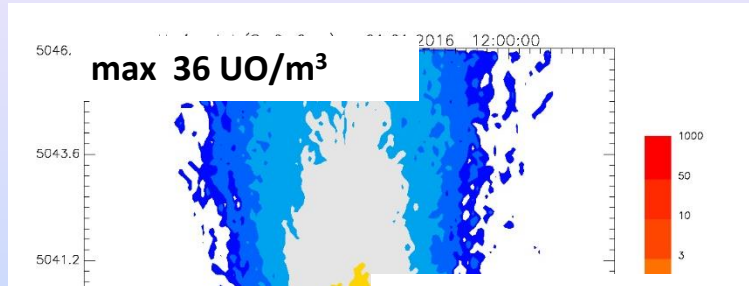
Unstable conditions, 1 m/s, concentrations

Average concentration $\langle C \rangle$



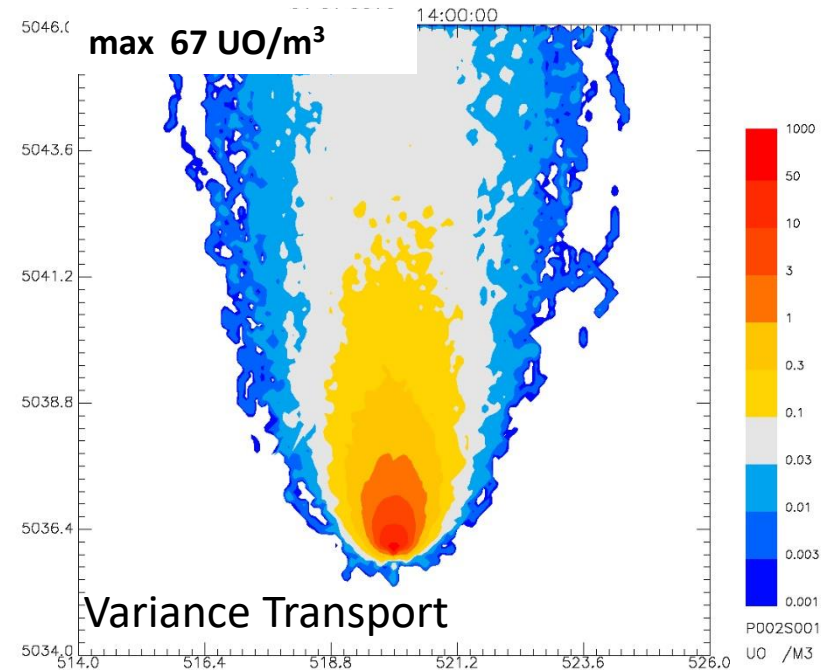
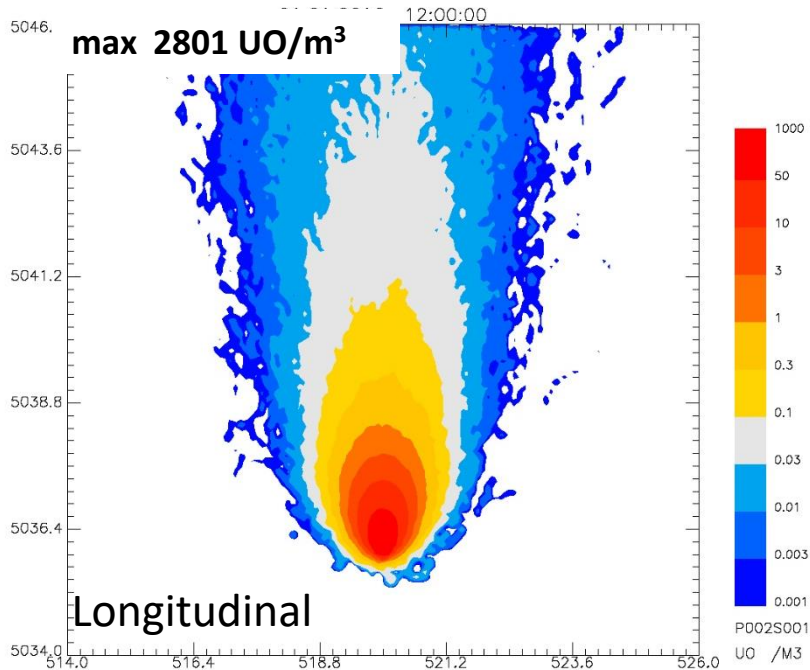
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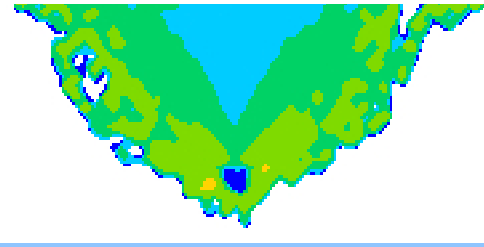
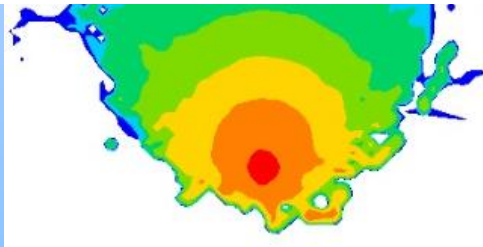
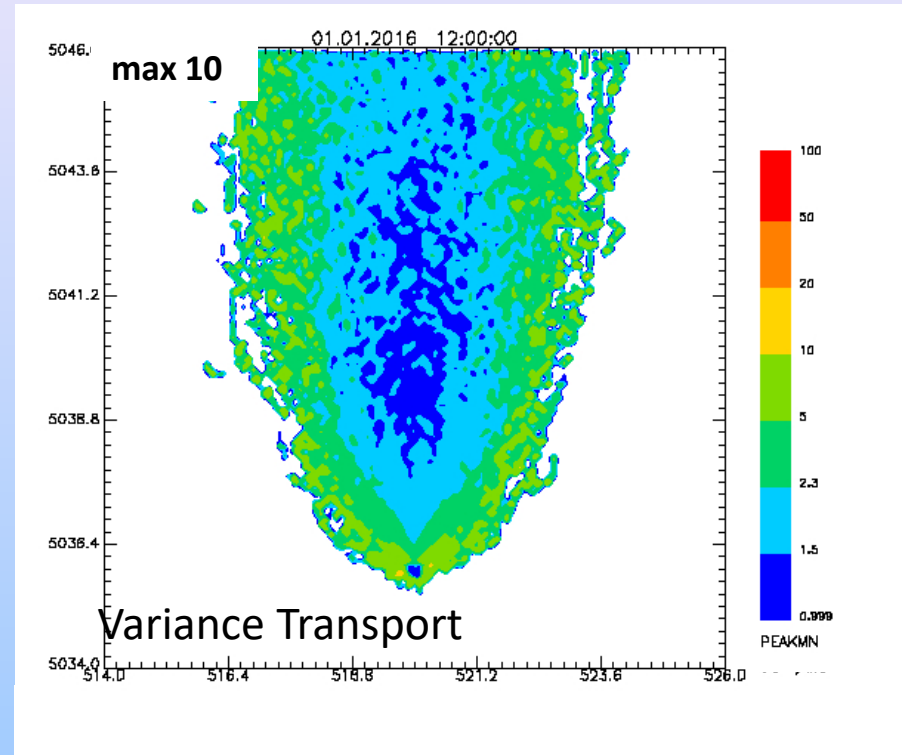
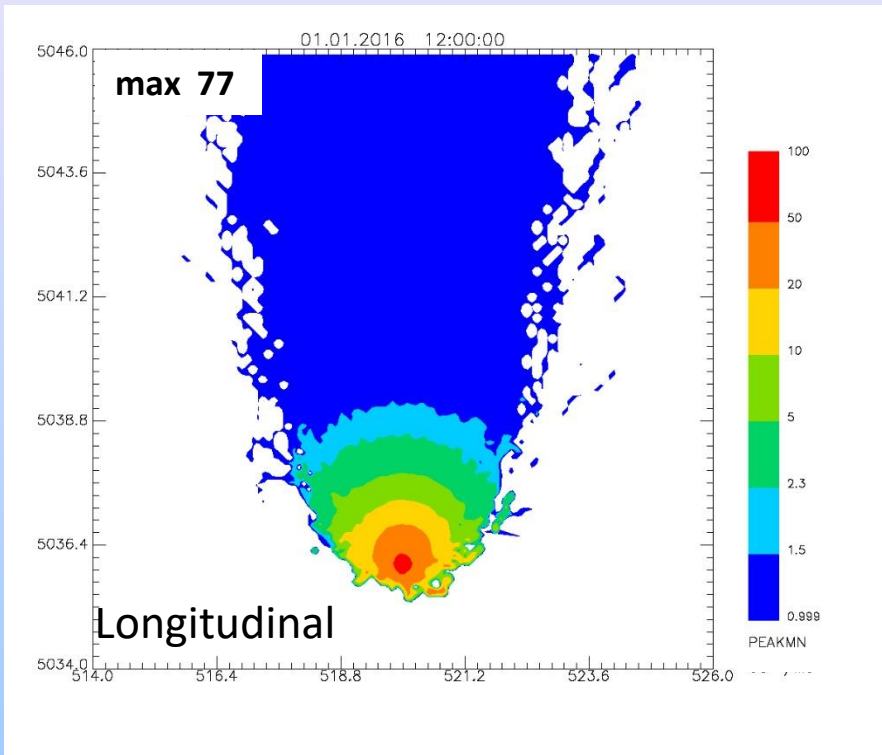


Peak concentration

Peak concentration

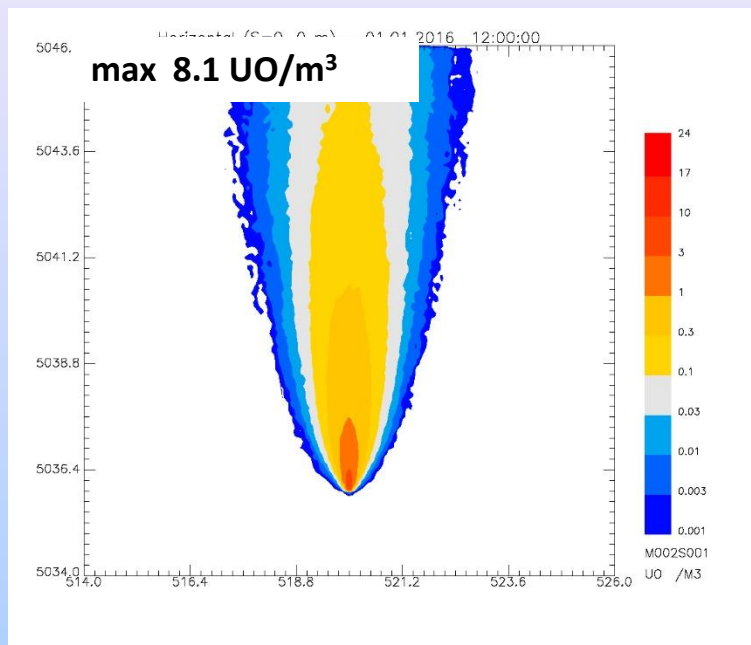


Unstable conditions, 1 m/s, Peak-to-mean ratio



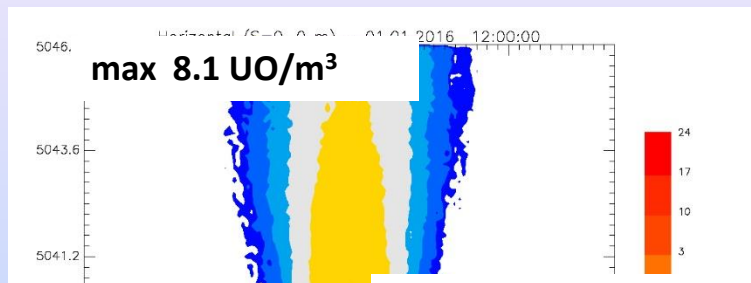
Unstable conditions, 4 m/s, concentrations

Average concentration $\langle C \rangle$



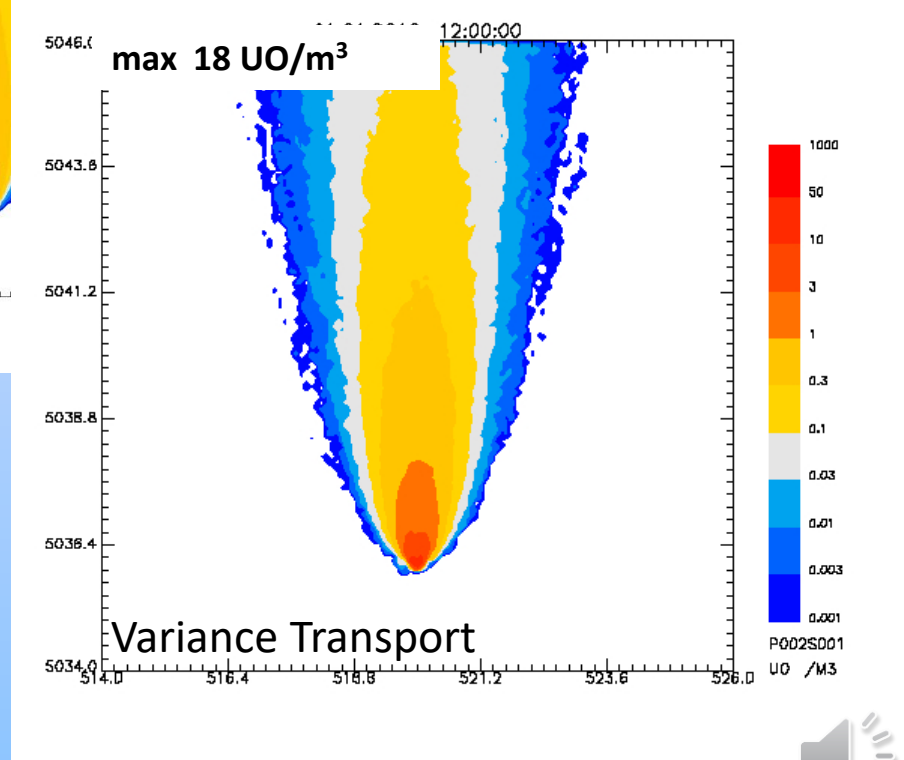
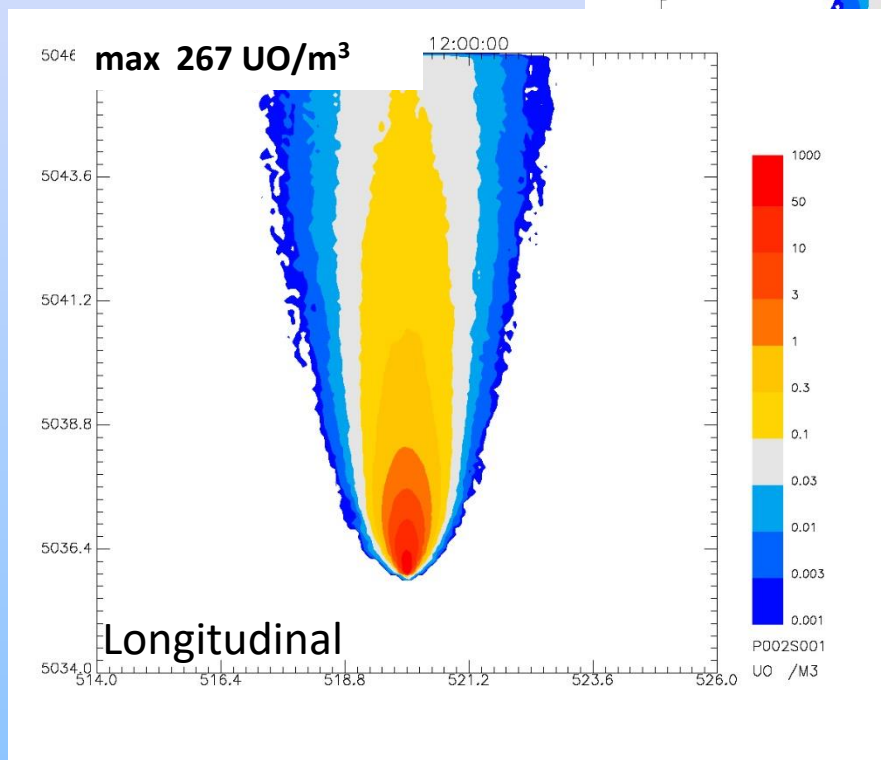
Unstable conditions, 4 m/s, concentrations

Average concentration $\langle C \rangle$

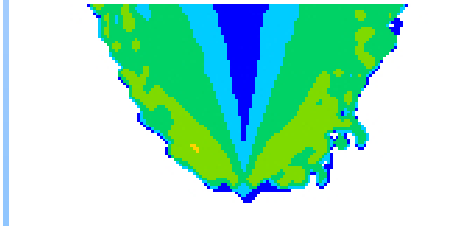
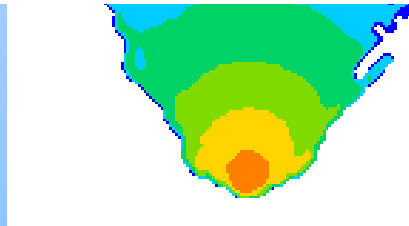
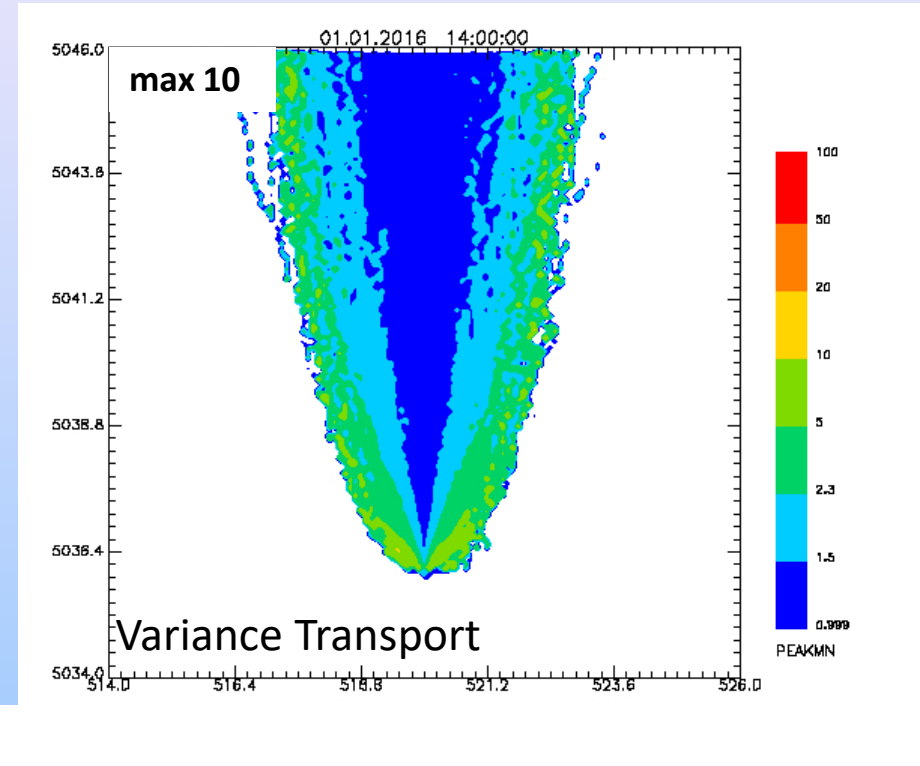
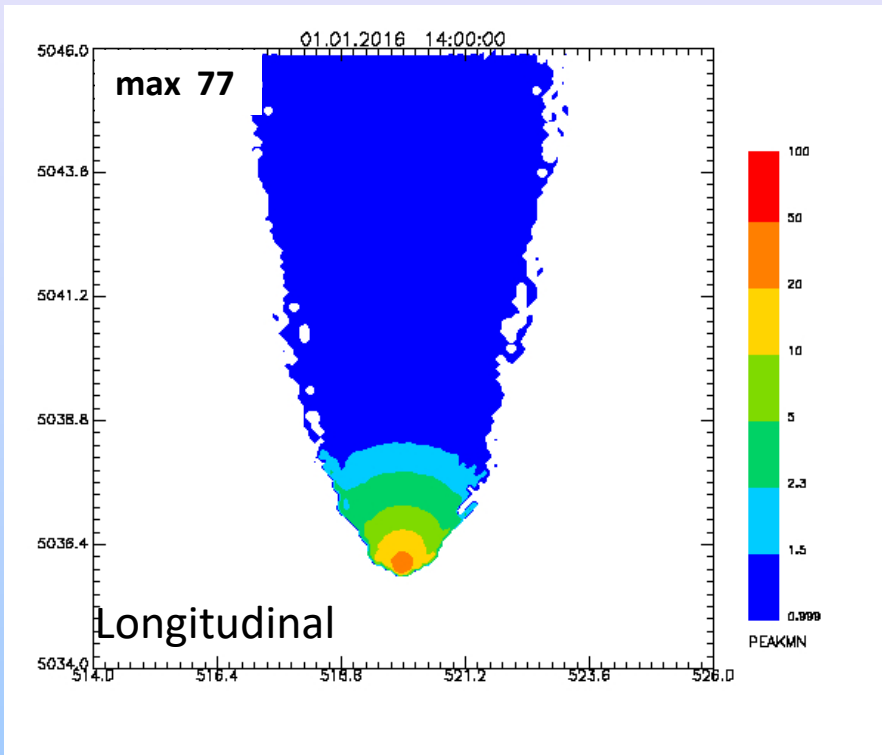


Peak concentration

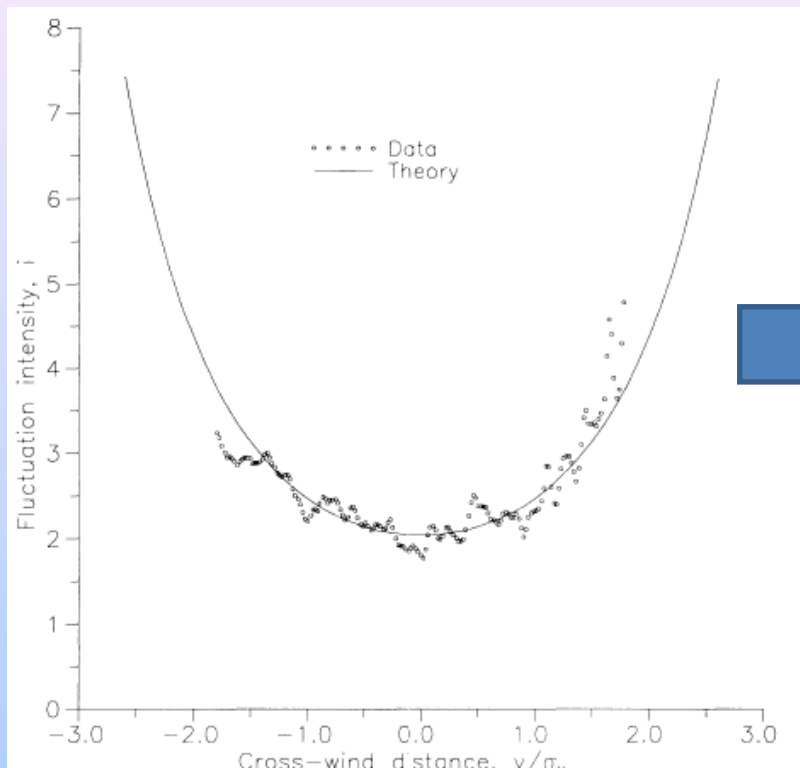
Peak concentration



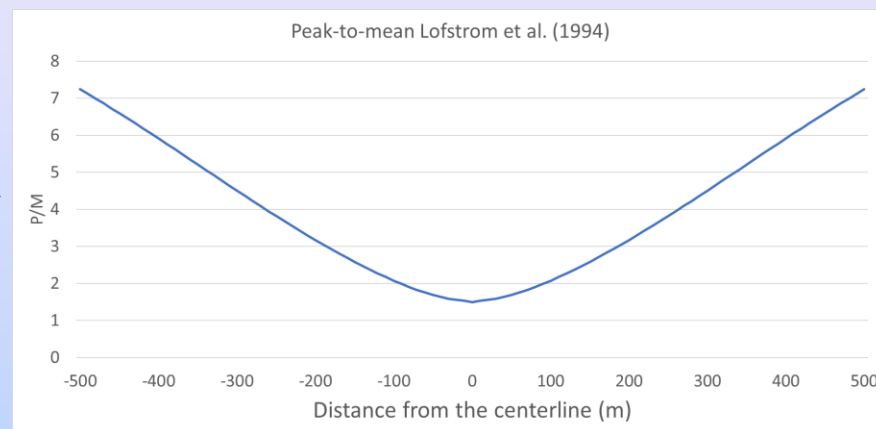
Unstable conditions, 4 m/s, Peak-to-mean ratio



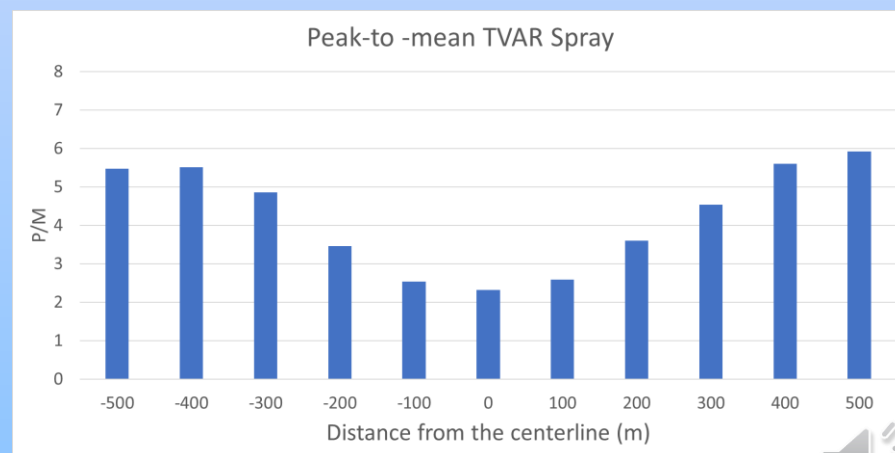
Lateral behavior of the fluctuation intensity, $\sigma_c / \langle C \rangle$, experimental data



Lateral Peak to mean from the interpolated curve
400m from the source,



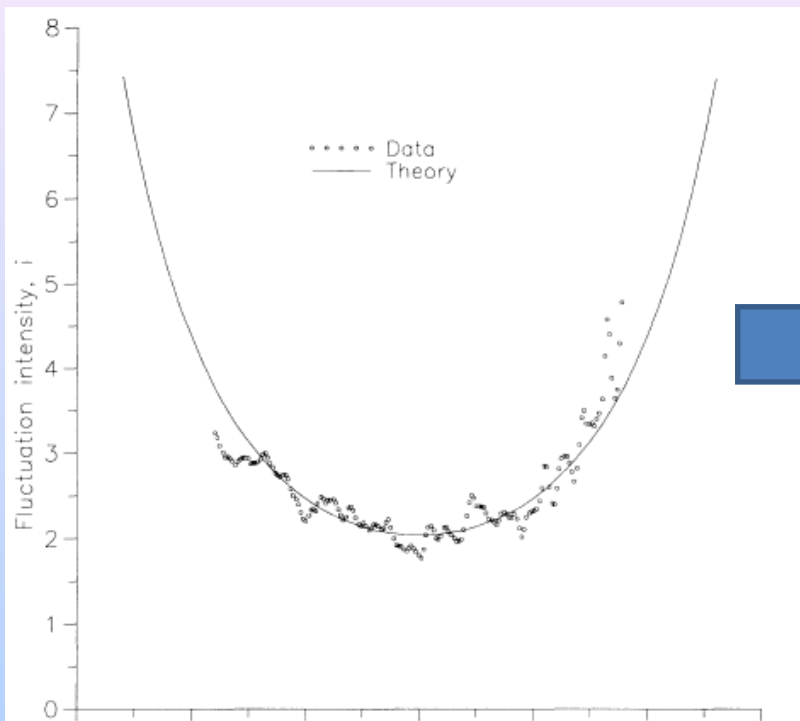
SPRAY 1 m/s unstable peak to mean



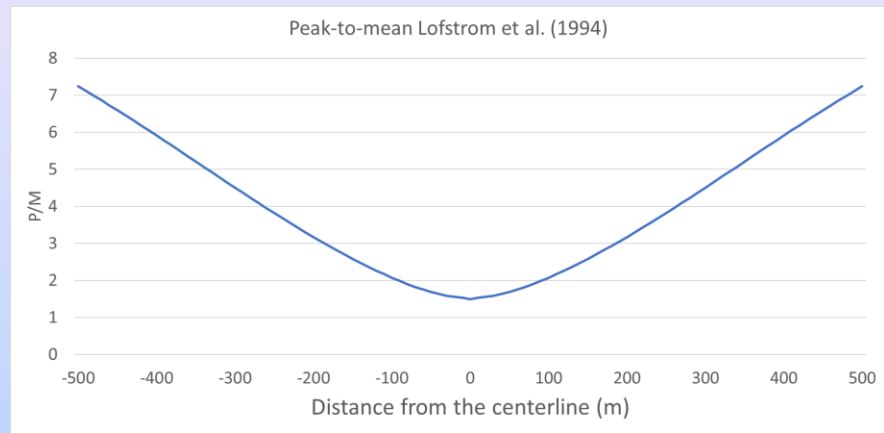
From Lofstrom et al., 1994

A concentration fluctuation model for decision-makers based on joint tracer and lidar measurements from a non-buoyant elevated plume – Trans. Ecology and Environ., 3, 571-579

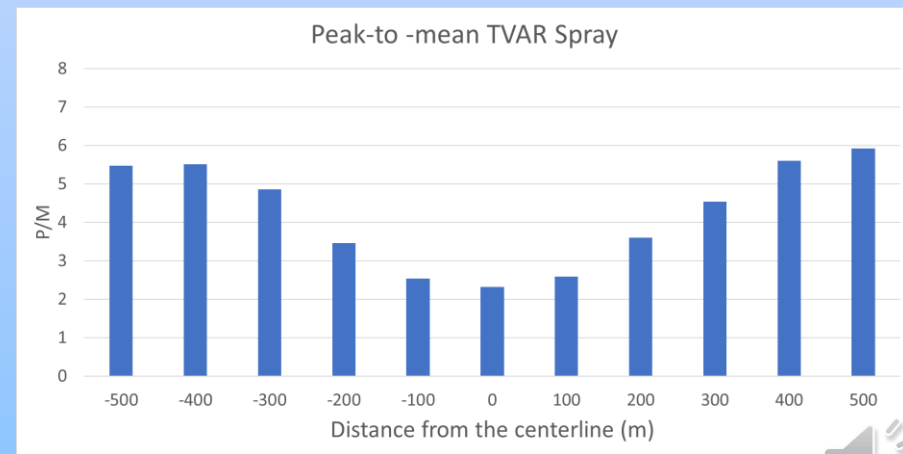
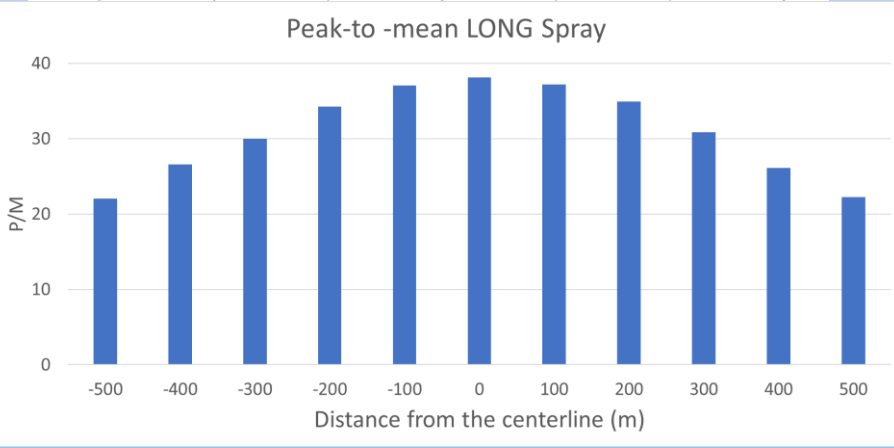
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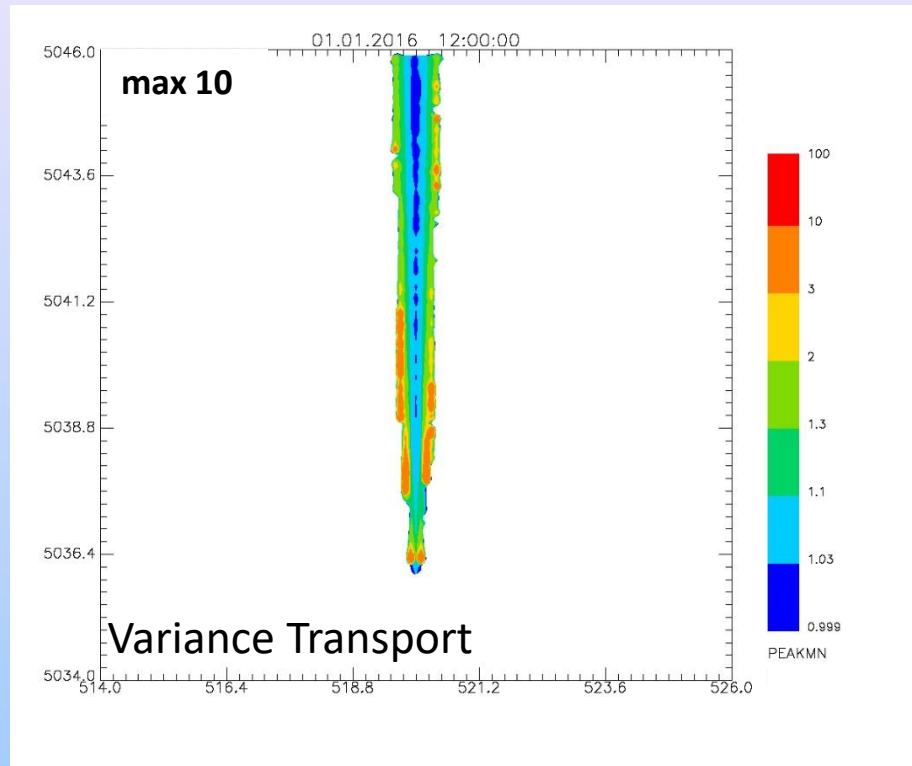
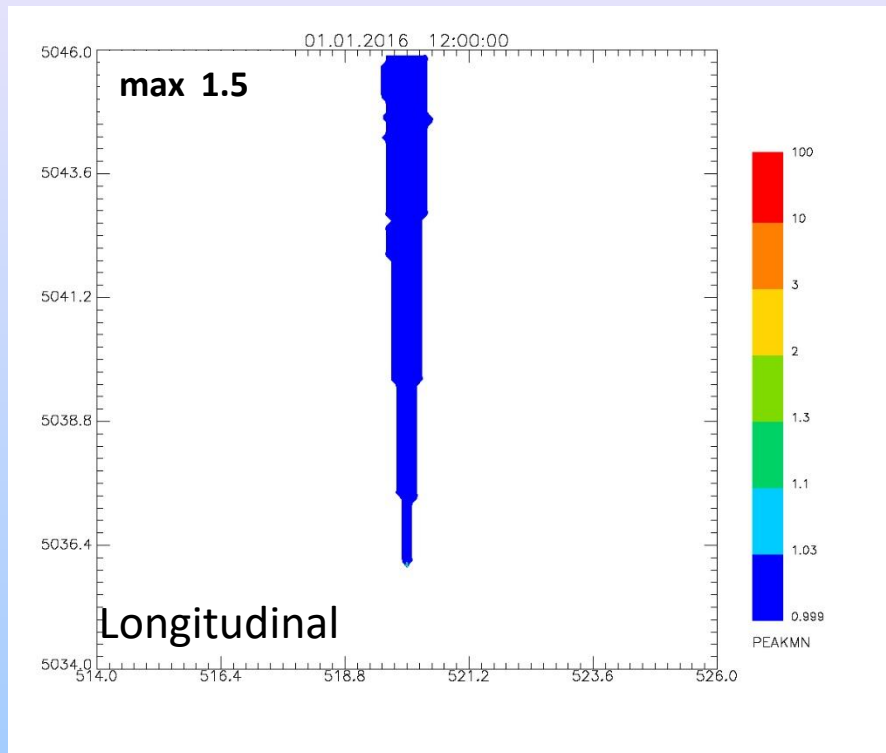
SPRAY 1 m/s unstable peak to mean



3. Numerical experiments - results

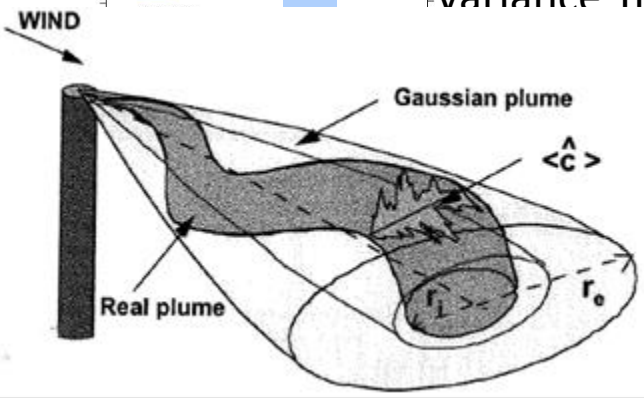
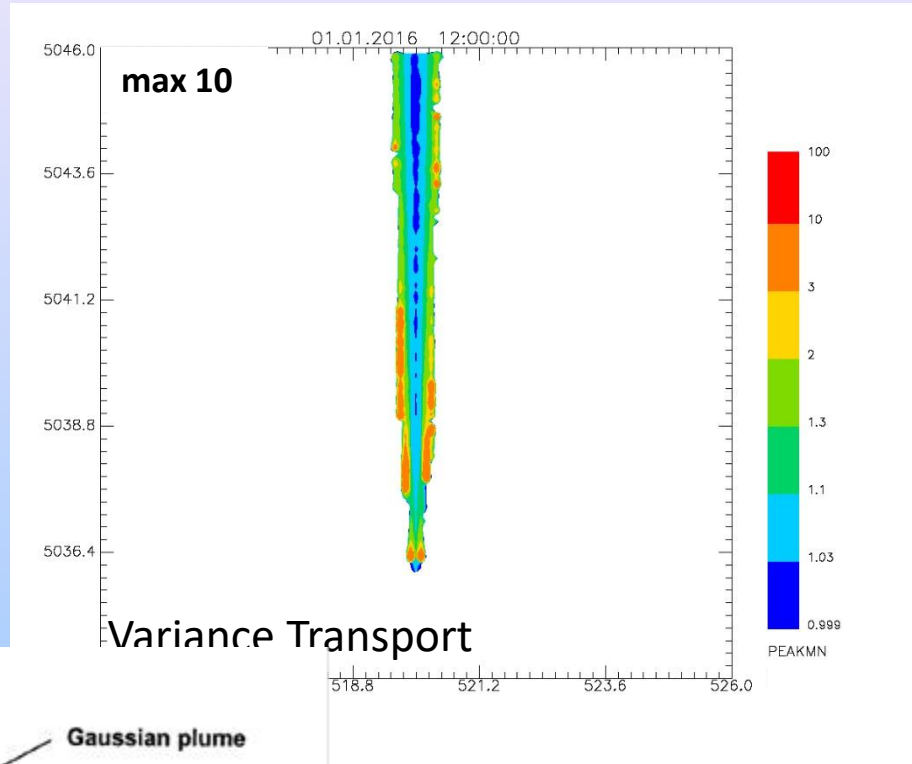
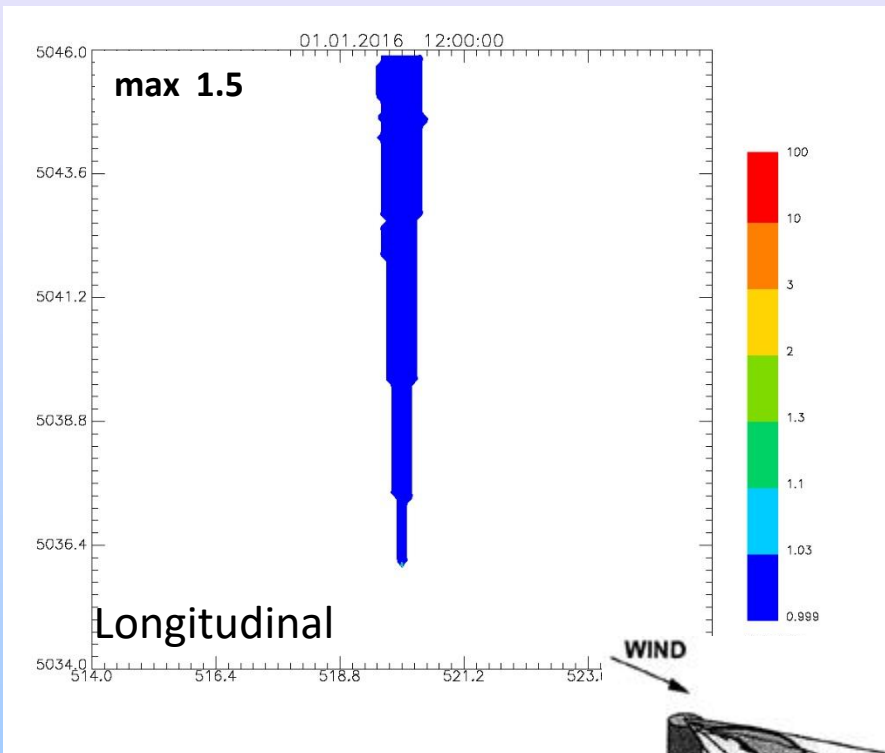


Stable conditions, 1 m/s, Peak-to-mean ratio

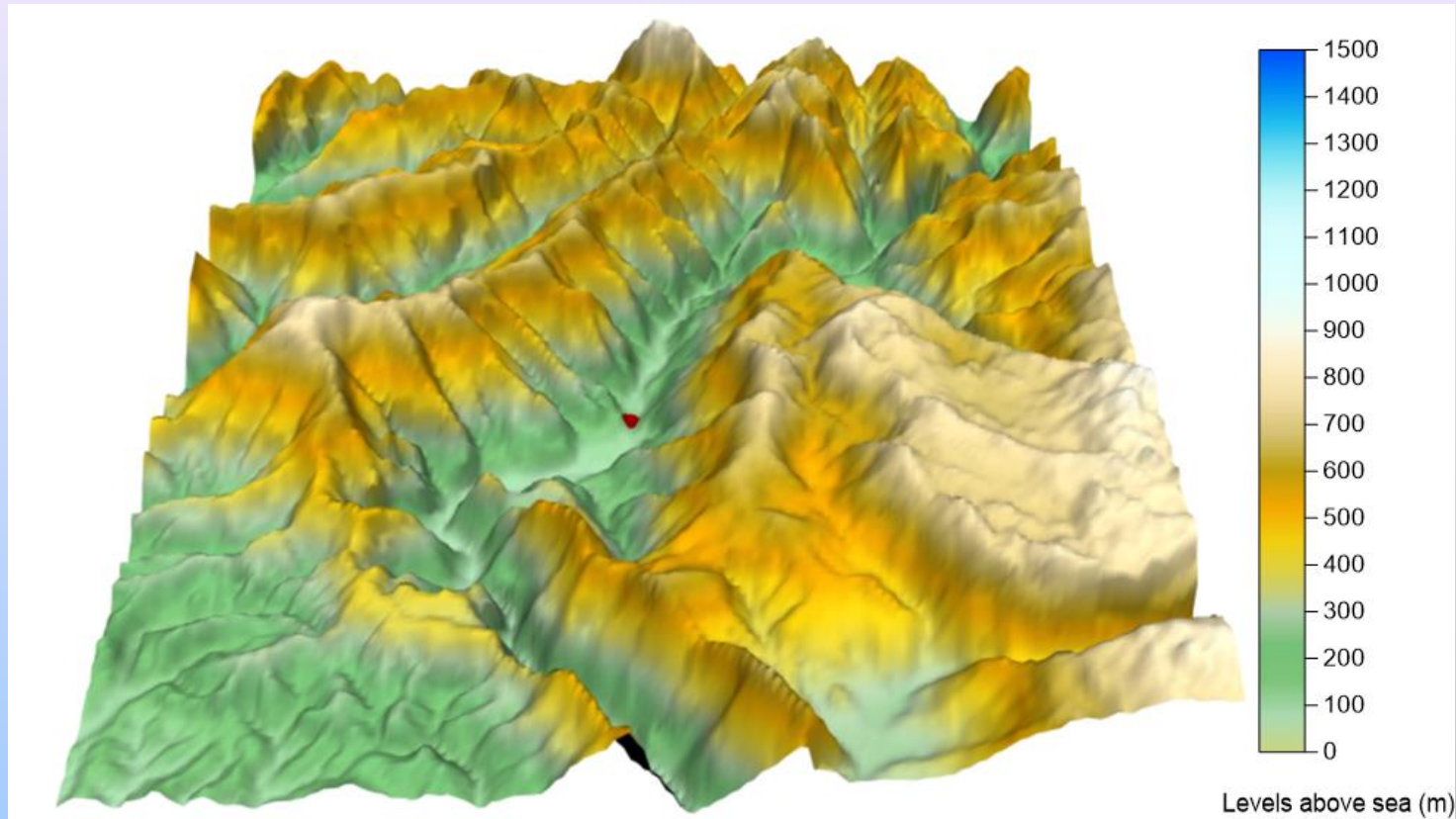


3. Numerical experiments - results

Stable conditions, 1 m/s, Peak-to-mean ratio



Realistic case over complex terrain



- 16 x 18 km² area
- Same emissions previously described
- 10 days during winter and 10 days during summer

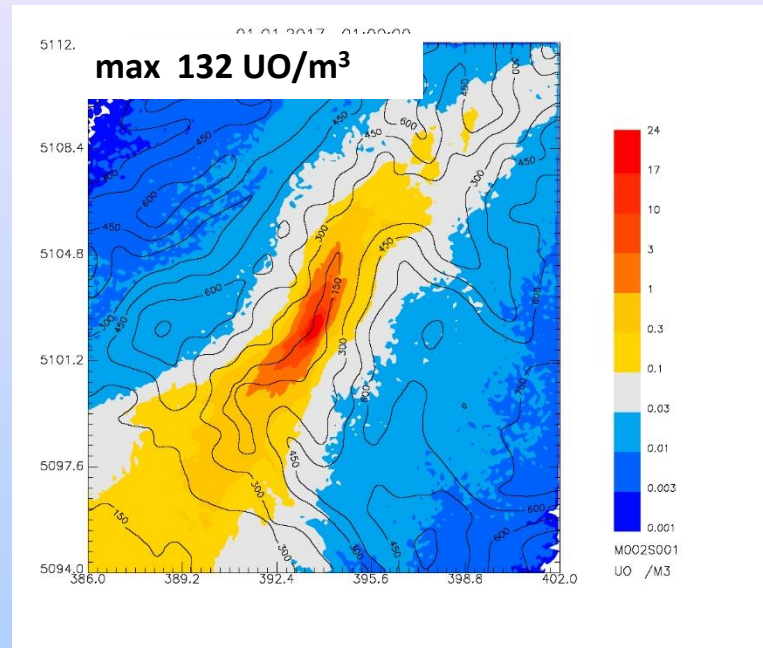


3. Numerical experiments - results



98 percentile derived from hourly averaged concentrations

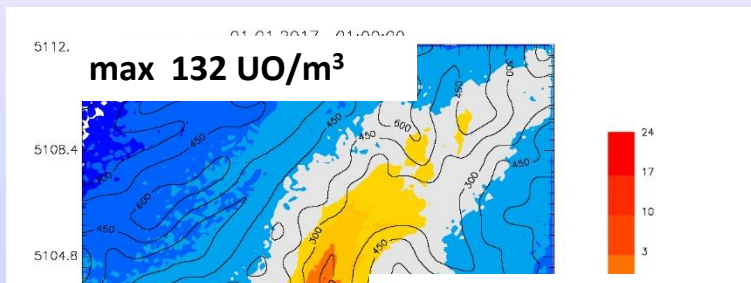
98 percentile <C>



3. Numerical experiments - results

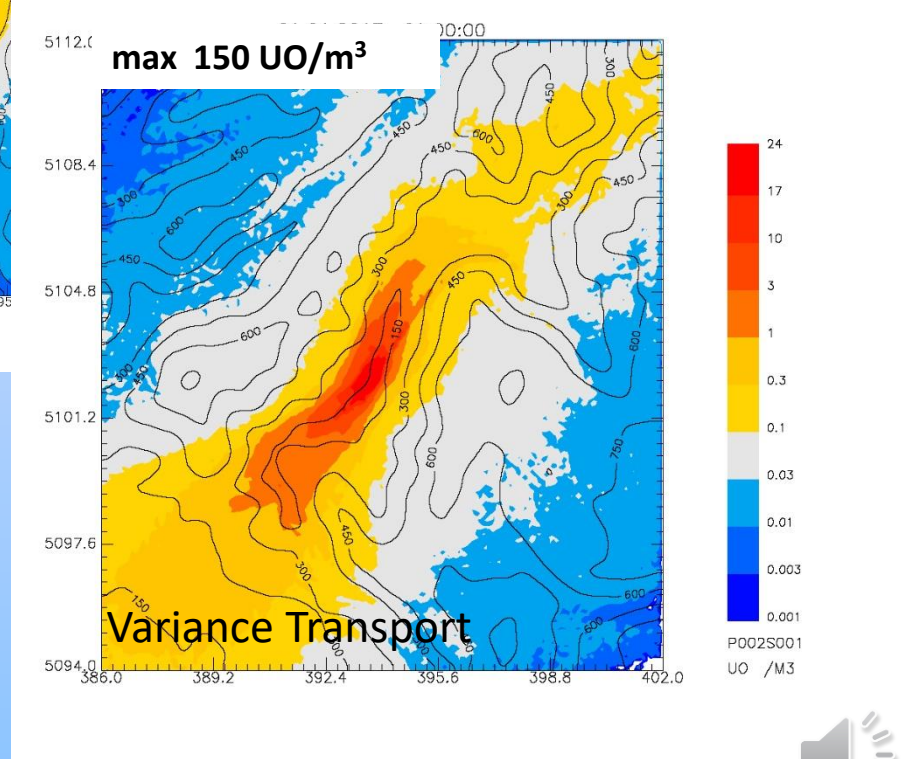
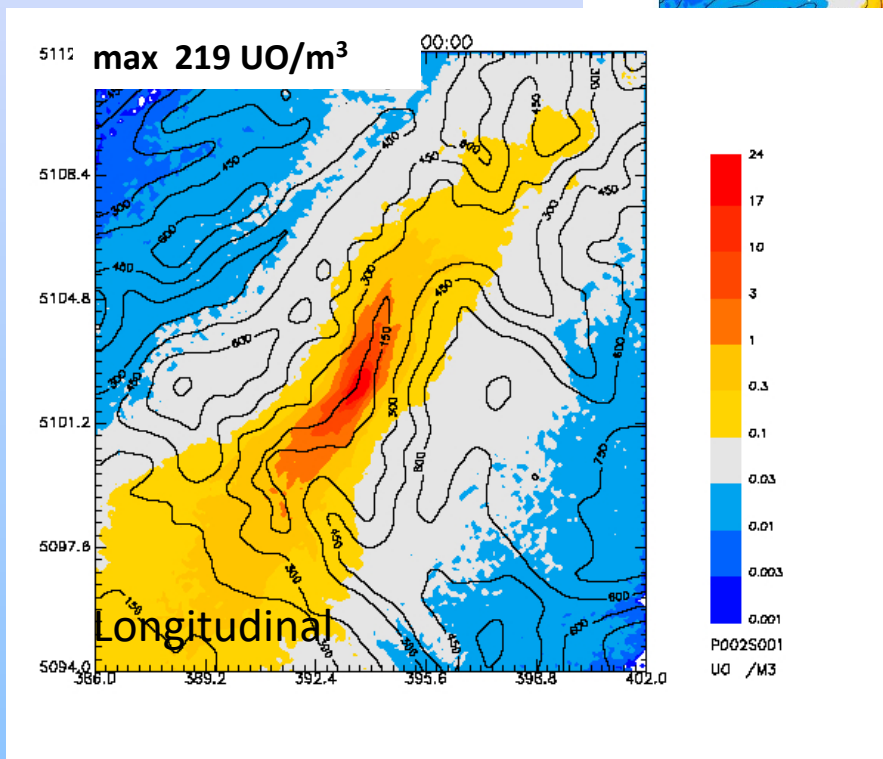
98 percentile derived from hourly averaged concentrations

98 percentile $\langle C \rangle$

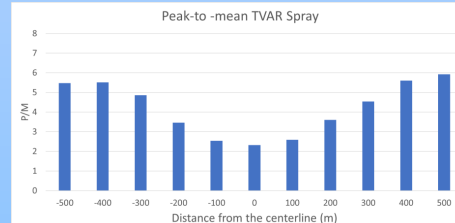
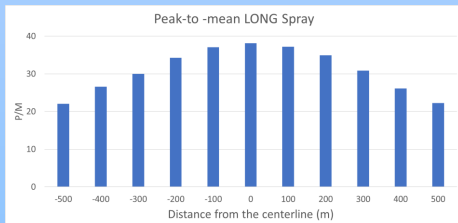
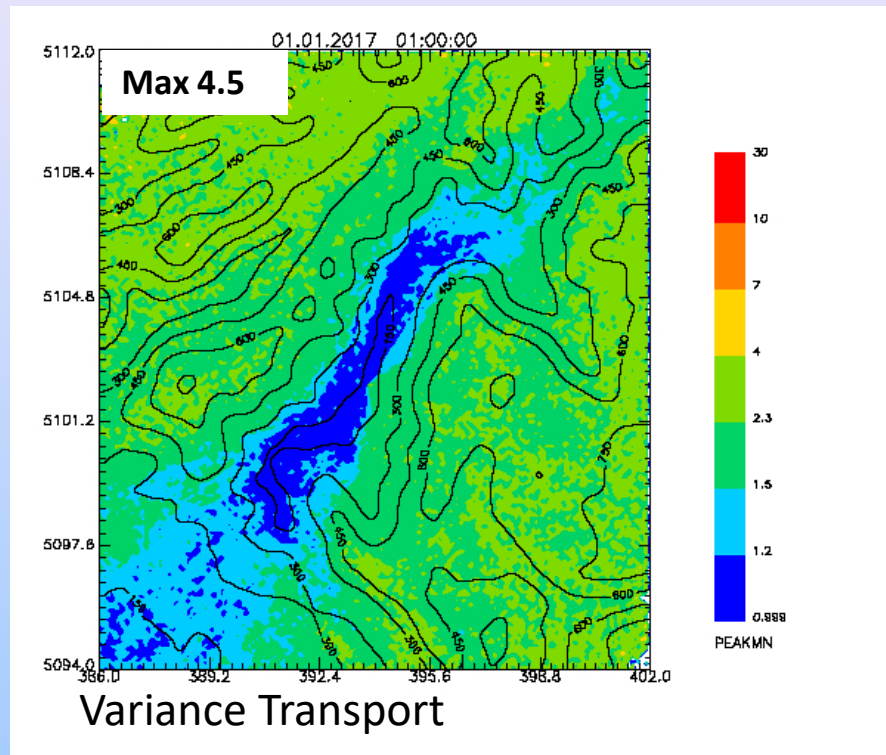
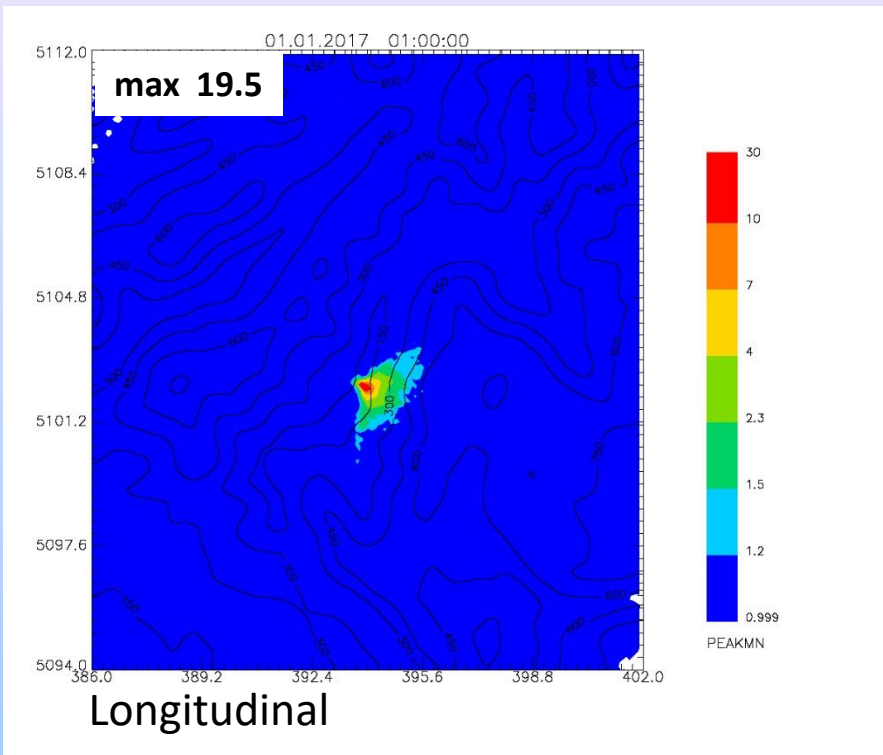


98 percentile Peak

98 percentile Peak



Peak-to-mean ratio of the 99.8 percentile



Some concluding remarks

- two P/M different approaches, having different spatial behaviours
- VART P/M seems to better follow some experimental evidences compared to the LONG approach, moving in a direction transverse to the plume
- stable cases: both approaches do not solve the problem of meandering to capture realistic peak values
- VART is oriented to a σ_c estimation, it has still some degrees of freedom moving to P/M (choice of the distribution and of the peak threshold)
- VART has the advantage to take into account multiple sources (it mainly depends on the concentration structure)



Thanks for your attention

